

What is sand?

- Where does it come from?
- Waves and wave energy
- How global climate change will influence waves and wave transport
- Dune types and formation**
- Natural dune communities
- Introduced beachgrass and influence on beach and dunes
- Surf zone and sand dwelling organisms and food web
- Snowy plover biology
- Oregon beach law
- Field trip

Dune types

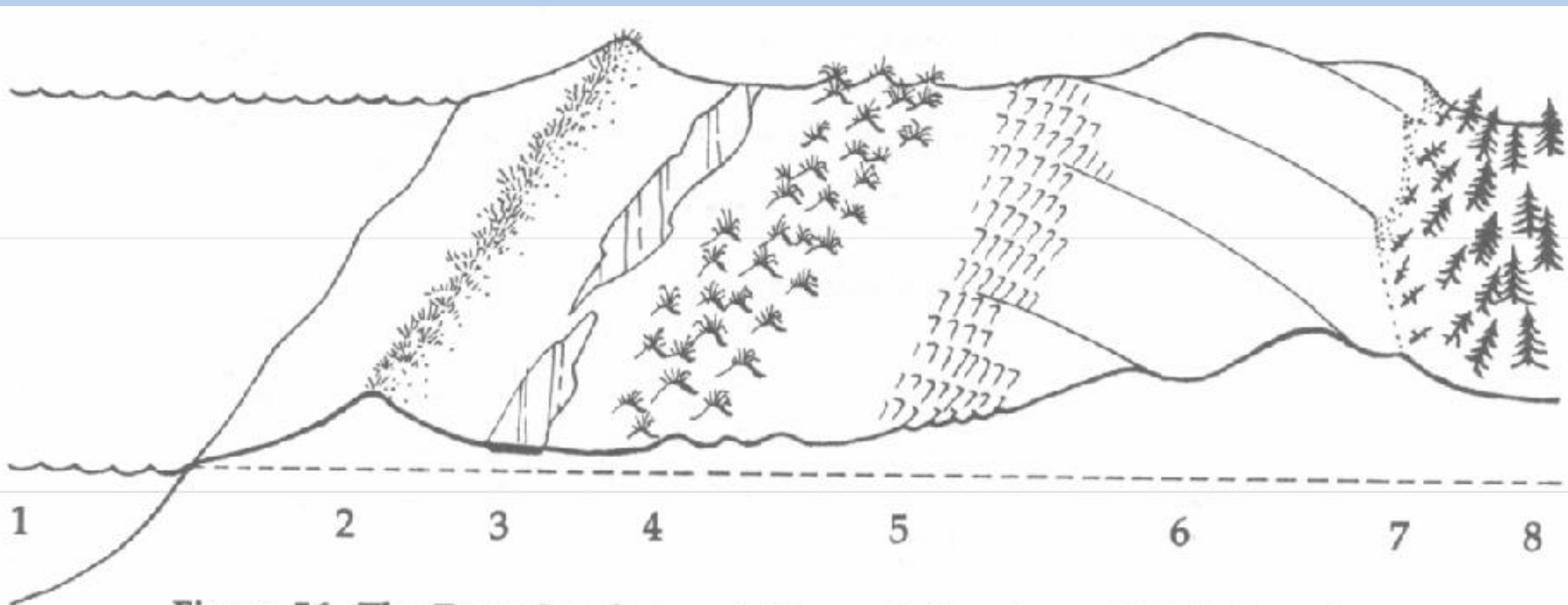


Figure 56. The Dune Landscape. 1 Ocean. 2 Foredune. 3 Deflation plain. 4 Beachgrass hummocks. 5 Transverse ridges. 6 Oblique dunes. 7 Retention ridge. 8 Forest. (Kellerman.)



Foredune

Pickart, 2008



Fig. 1. The native dune grasses *Leymus mollis* and *Poa macrantha* mix with forbs, including *Abronia latifolia* (yellow sand verbena) and *Lathyrus littoralis* (beach pea), on the foredune at the Lanphere Dunes Unit, Humboldt Bay National Wildlife Refuge.



Fig. 6b. The same location in July 2001, five years after restoration work was completed.

Pickart, 2008



Fig. 10. Restored foredune grassland at the Lanphere Dunes Unit, Humboldt Bay National Wildlife Refuge.

Pickart, 2008

Deflation plain



Beachgrass hummocks



Transverse dune



Dune trend 80°

Wind direction
and dune migration
170°



Oblique dune





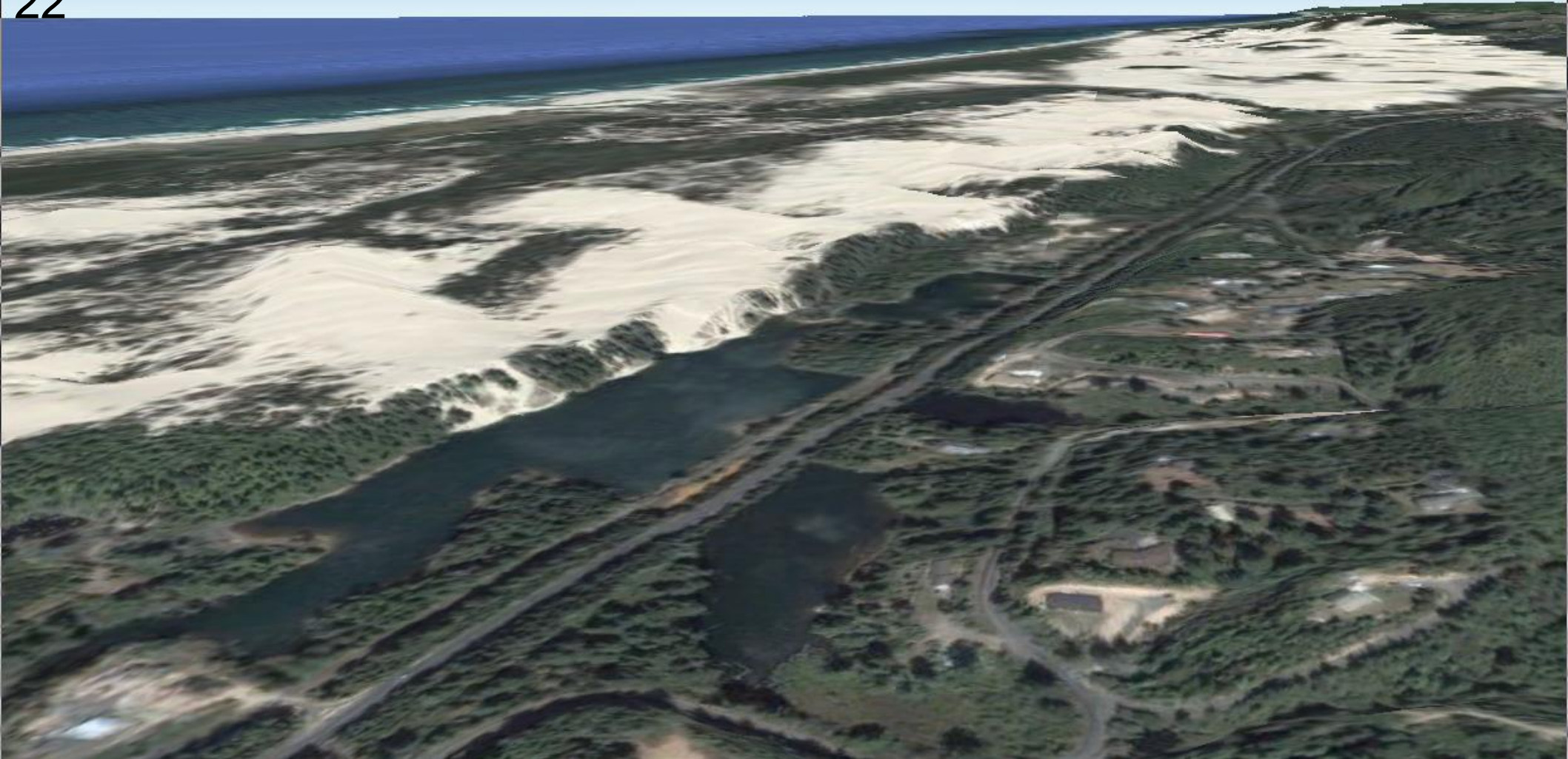




Dune advance 26°
Resultant sand transport 45°
Resultant winter wind
trend 116°

Retention ridge





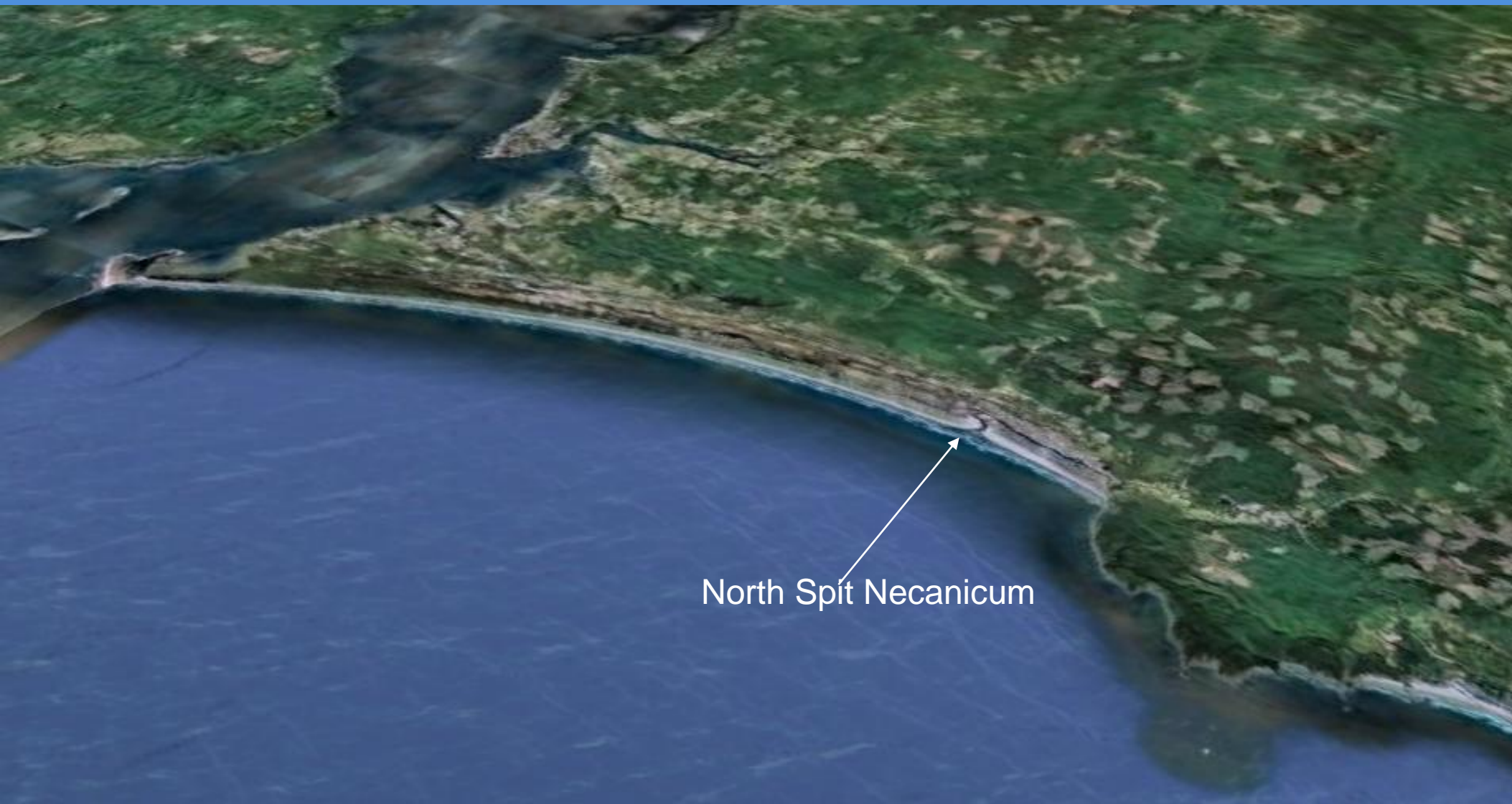
Parabola dune







Parallel ridges



(1) Columbia River south jetty to Tillamook Head; 26 km

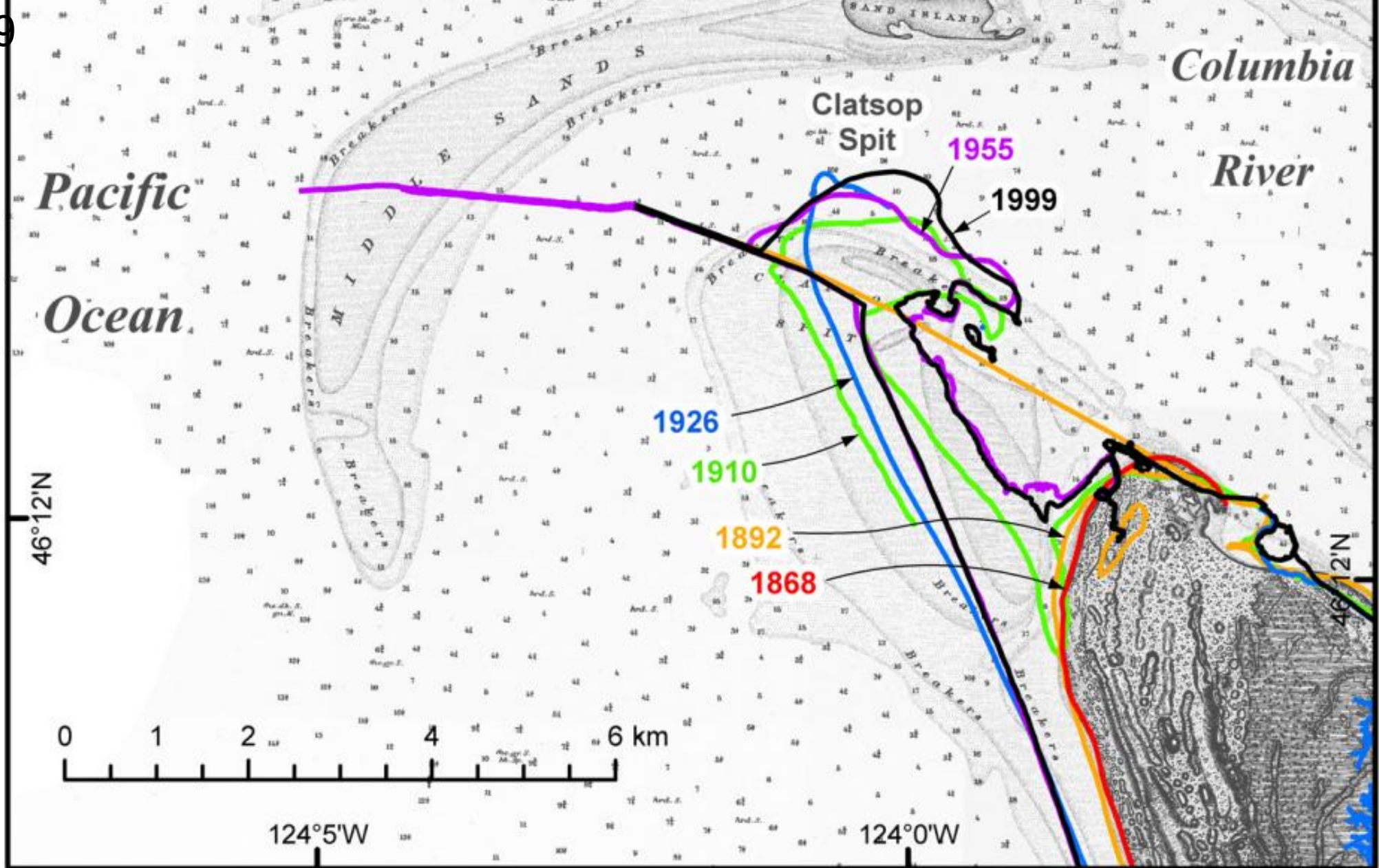


Fig. 6. 1870 U.S. Coast and Geodetic Survey – Mouth of the Columbia River, with historical shorelines.



Cooper's paradigm of dune development

Figure 55. Sea level rise and dune advance.
(Kellerman.)

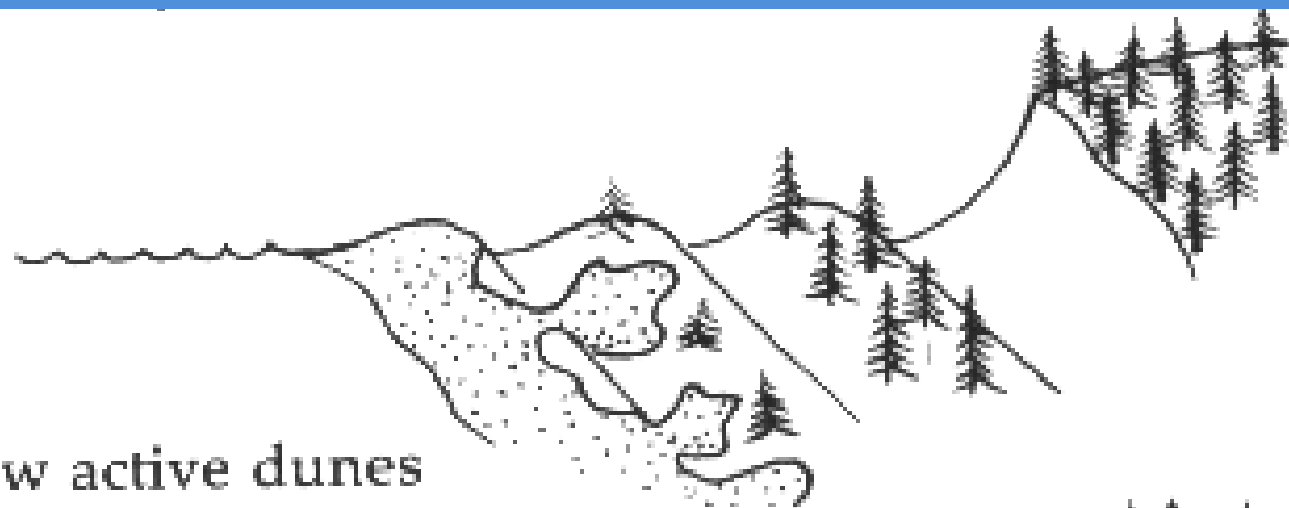


Previous dune landscape is stabilized during quiescent period.



Sea level rises, re-activating old stabilized dunes.



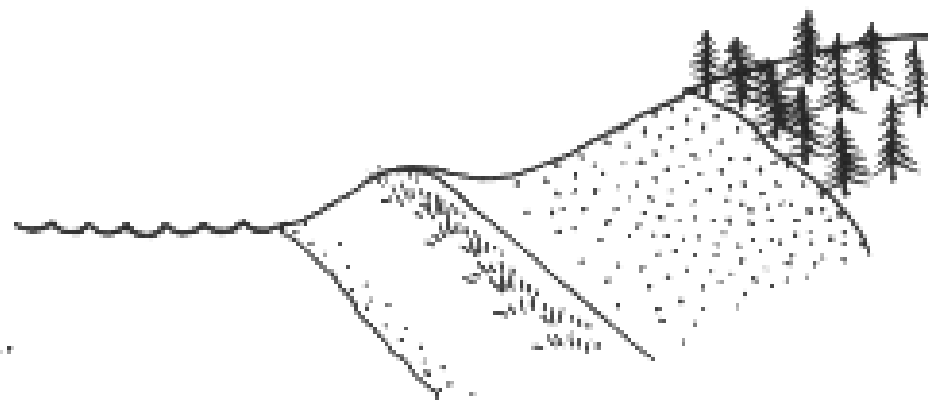


Sea continues to rise, and new active dunes spread inland.

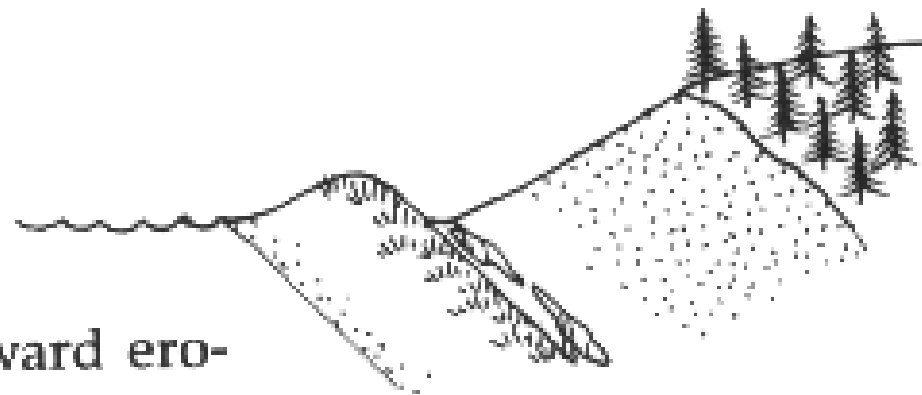


Active dunes reach maximum advance, forming oblique ridges.

Sea level stabilizes, foredune is created.



Foredune traps beach sand, allows leeward erosion to water table, creating deflation areas.

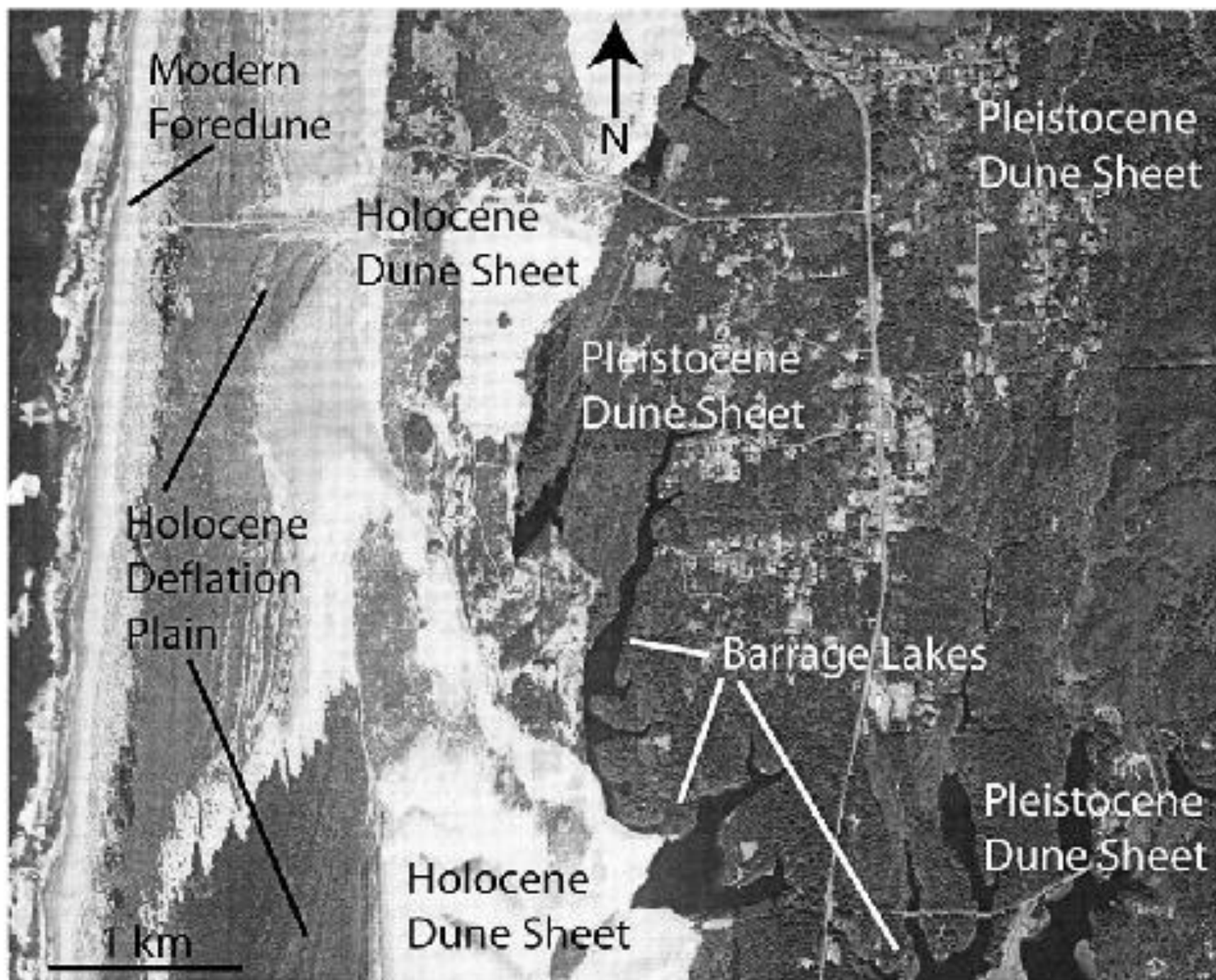


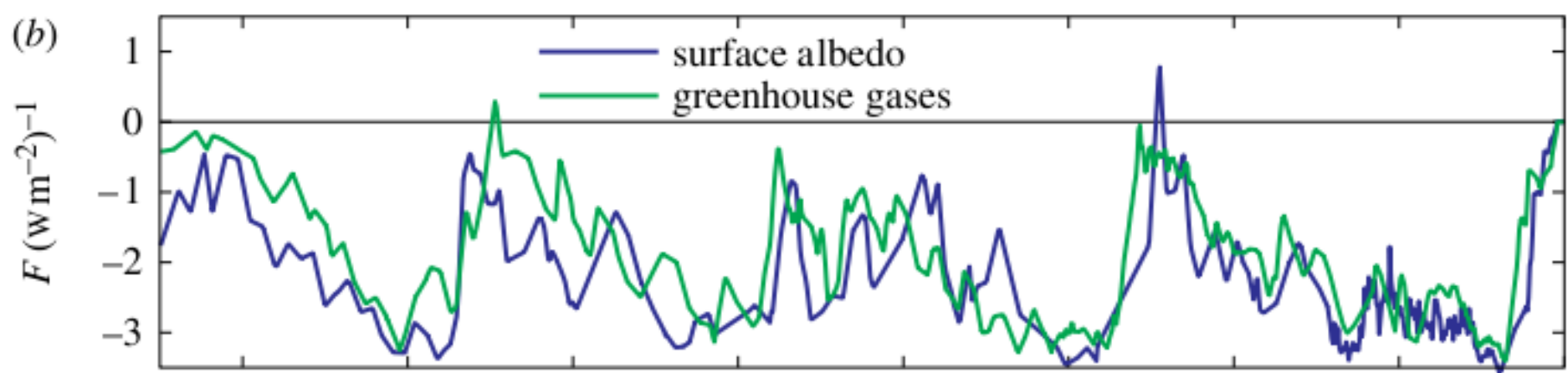
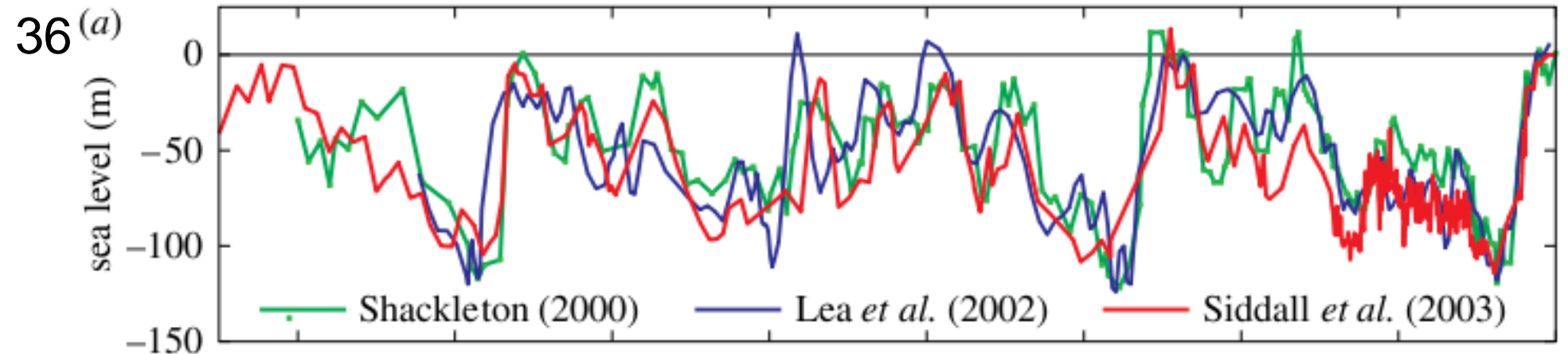
Deflation areas colonized by vegetation.



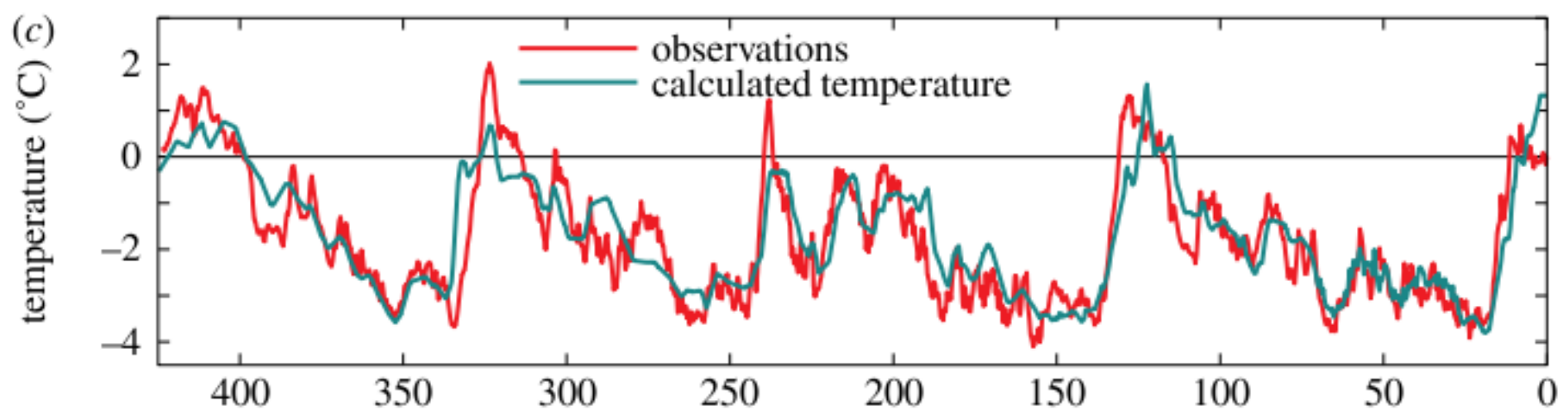
Eventually previous active sand is stabilized as succession progresses to forest.





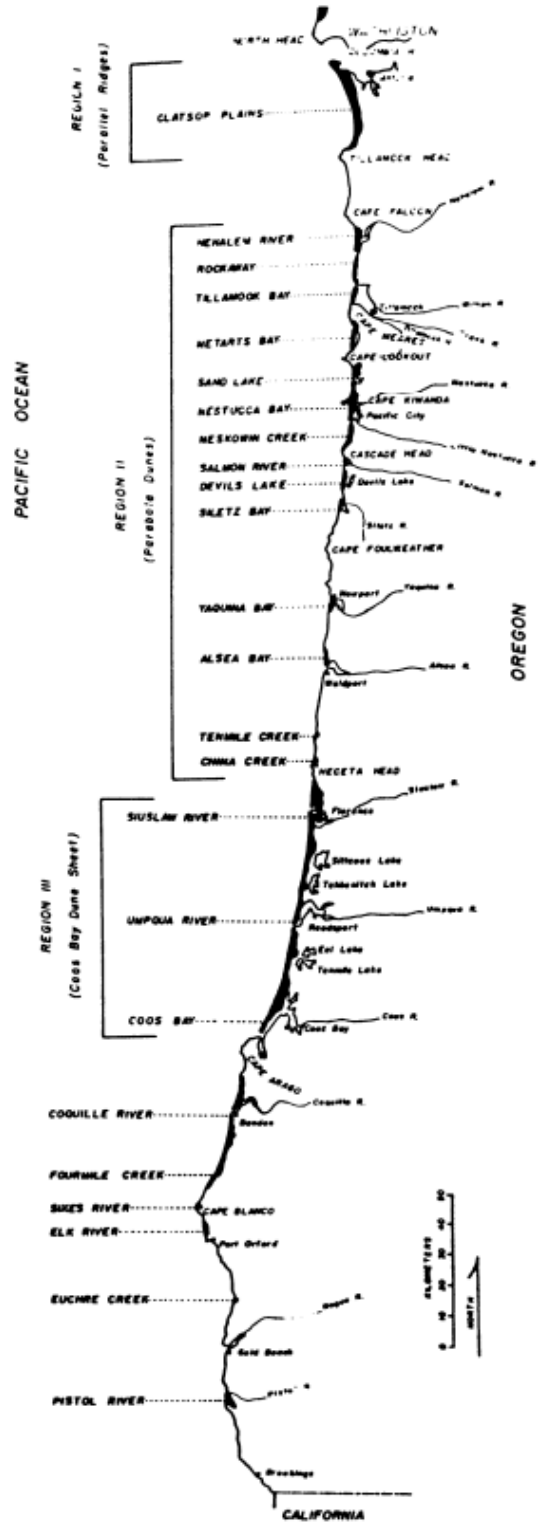


Hansen, et al., 2007

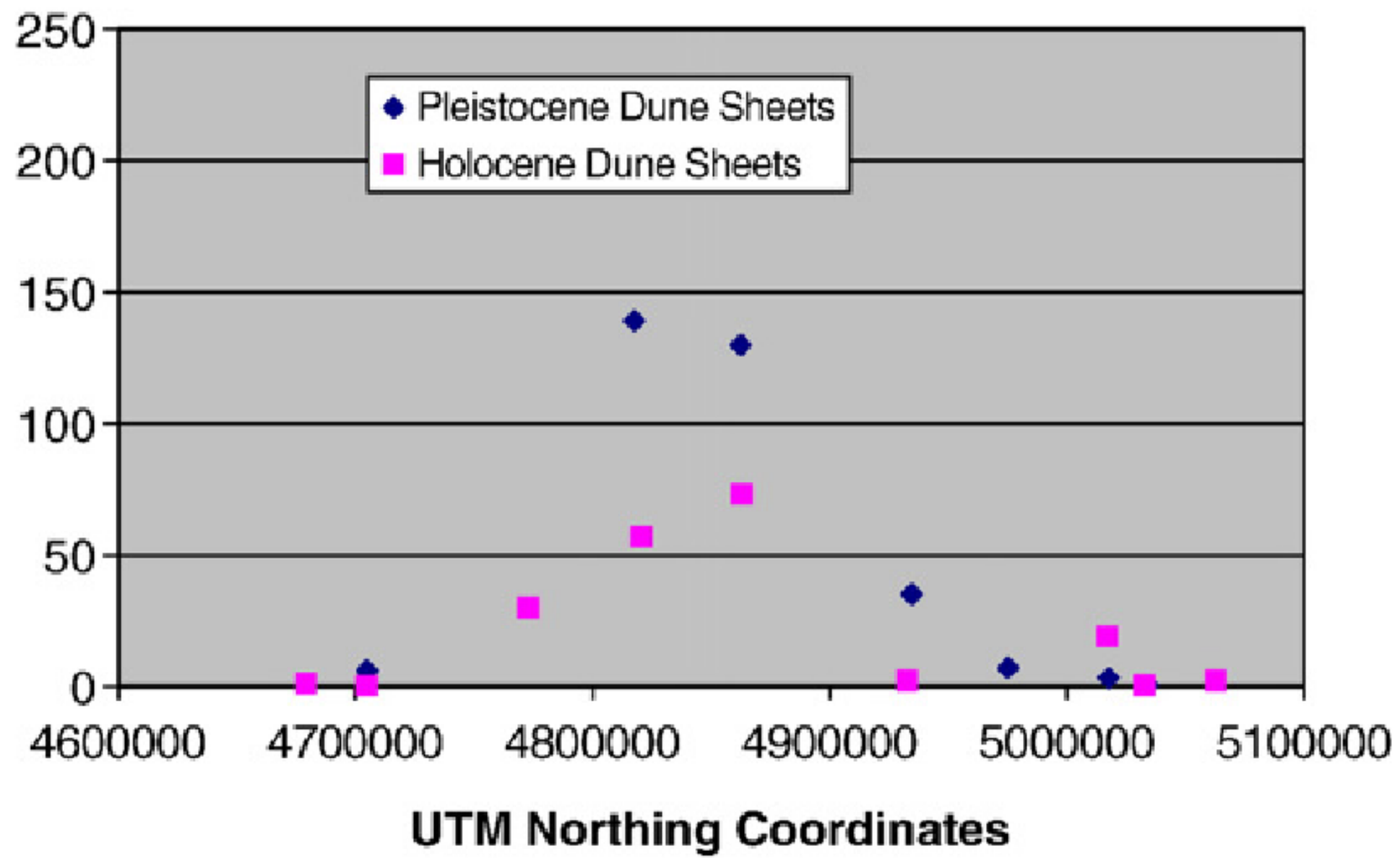


PACIFIC OCEAN

OREGON



Dune Sheet Size versus Position



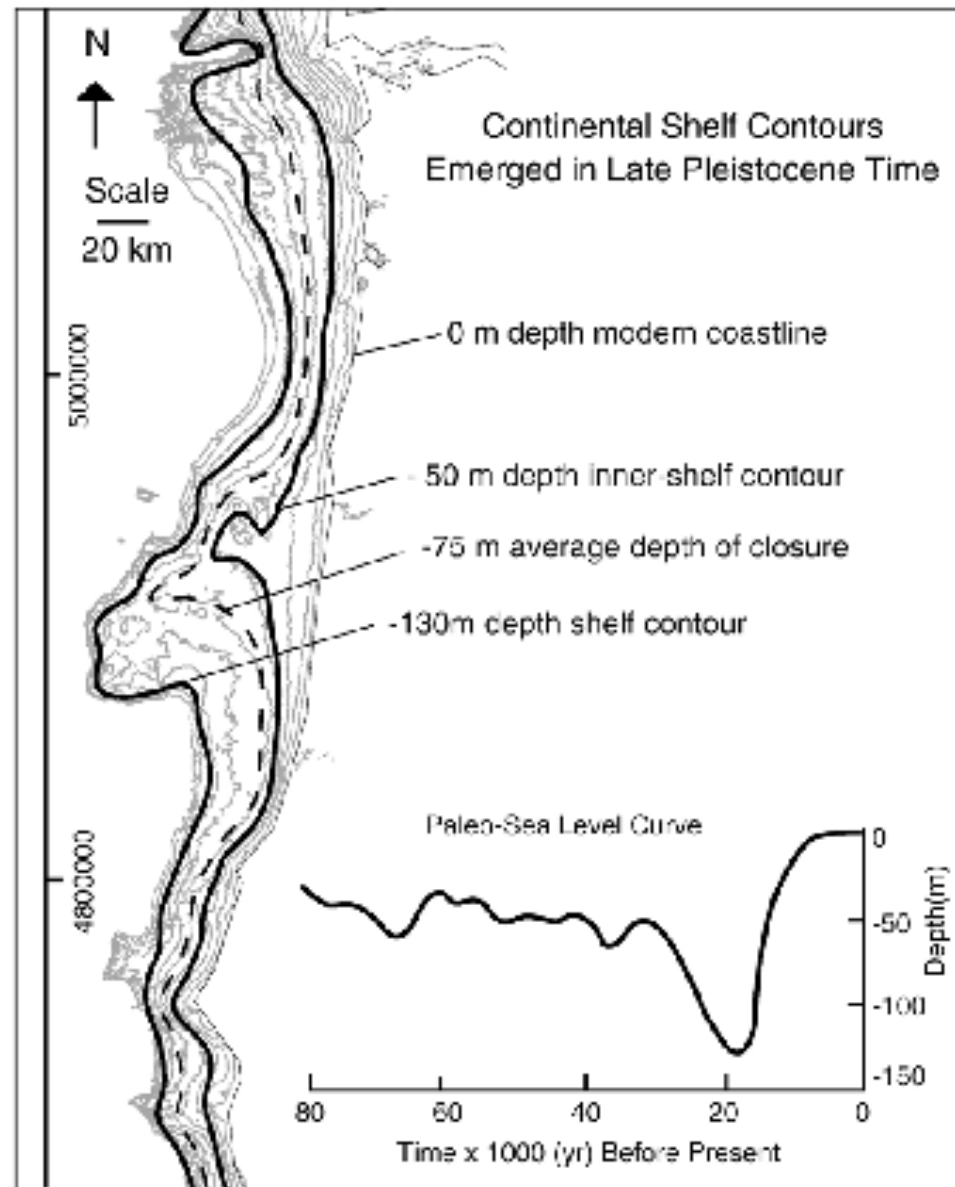
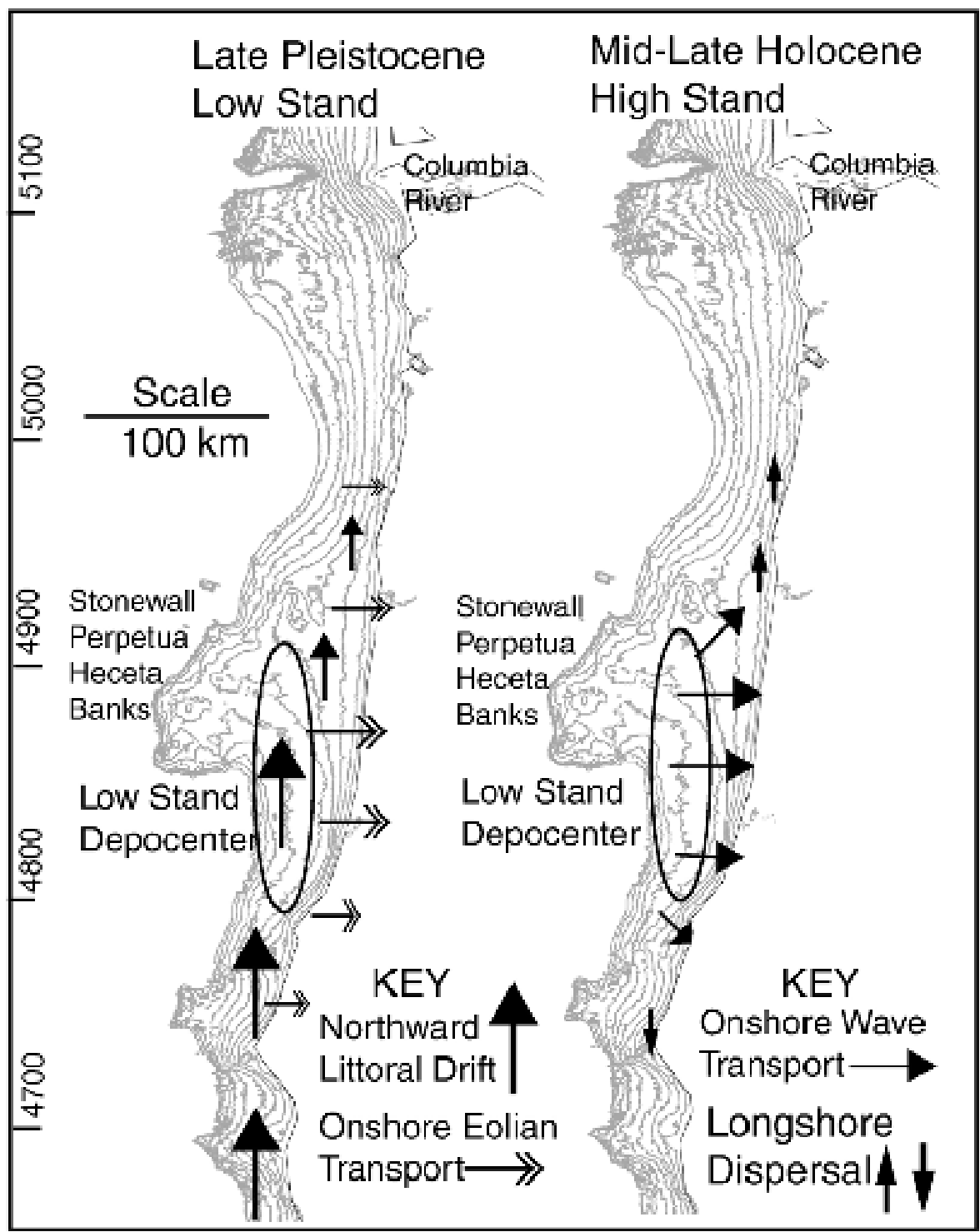


Fig. 4. Map of Oregon continental shelf showing paleo-sea levels (solid lines) at -50 m and -130 m isobaths. A eustatic, sea-level curve is shown for the last 80,000 years (figure redrafted from Pirazzoli, 1993, sea-level curve from Chappell and Shackleton, 1986). The major period of Late-Pleistocene dune emplacement (30–70 ka) is dominated by sea levels of about 50 m below present sea level. The average depth of closure, e.g., limit of littoral transport (dashed line at -75 m contour), for the late Pleistocene is assumed to have been about 25 m below the average sea level for that period. Sea level during the last glacial maximum (21–18 ka) briefly decreased to 130 m below present sea level.



Dune formation

- Most dunefields are from the Pleistocene, 50 kya
- Sea level not rapidly rising then, and sea at around 50 m depth line
- No dunes older than around 80 kya, not known why
- Climate was cooler 50 kya, and might have been drier, but not known. Maybe weaker stabilization?

References

Pickart, A.J. 2008. Restoring the grasslands of northern California's coastal dune. *Grasslands* 17(1): 3-9.

Peterson, C.D. et al. 2007. Ages, distributions, and origins of upland coastal dune sheets in Oregon, USA. *Geomorphology* 91(1-2): 80-102.